



Magmatic and **Current Deposit**

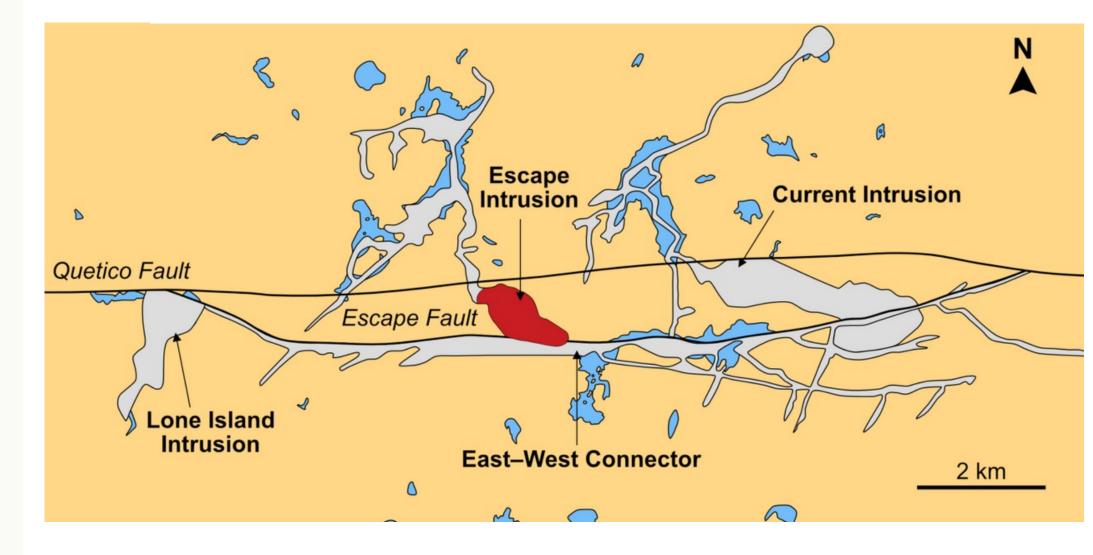
Corredor, A.P., Hollings, P., Brzozowski, M., and Heggie, G.



hydrothermal evolution of the PGE-Cu-Ni

Thunder Bay North Intrusive Complex

- Several mineralized and nonmineralized mafic-ultramafic intrusions [1] that intruded the Quetico Basin in the early stages of the of the MRS (Period of extension of the lithosphere with a massive magmatism; [2,3])
- Clean Air Metals is the current owner of the Thunder Bay North Project.



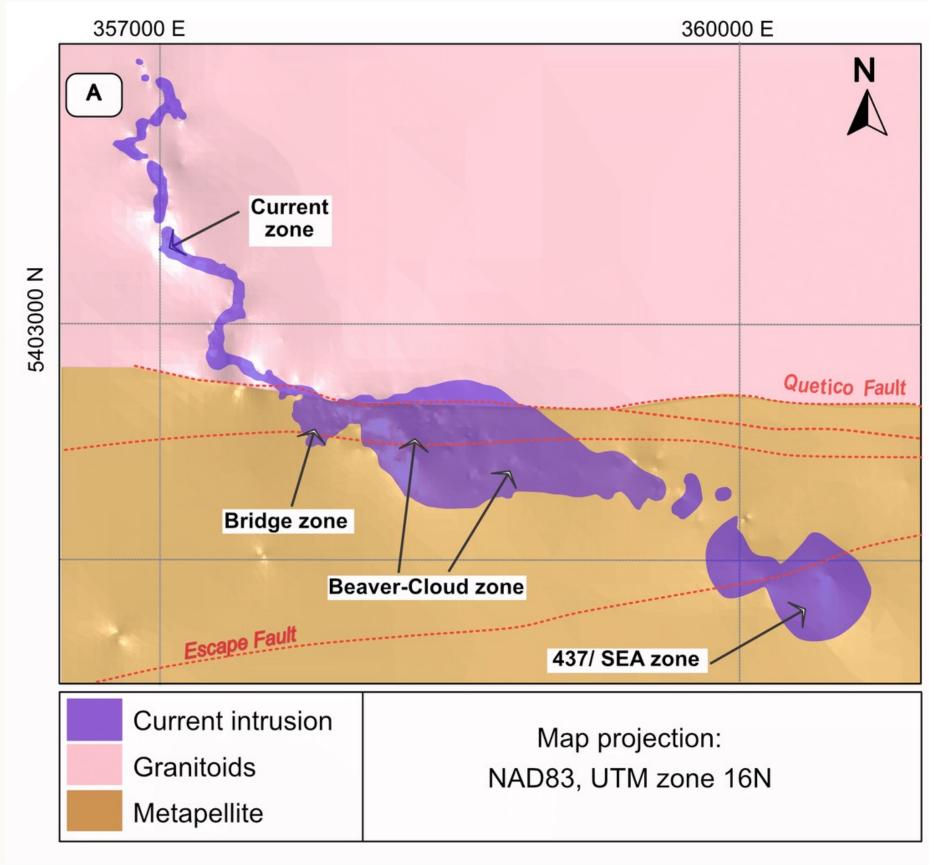
[1] Bleeker et al. (2020) [2] Woodruff et al. (2020) [3]Hinze & Chandler (2020) [4] Caglioti , 2022

Thunder Bay North Project map [4]

Current Intrusion

- The 1,106.6 ± 1.6 Ma intrusion
 hosts PGE-Cu-Ni mineralization

 [1].
- The intrusion is associated with multiple faults of the Quetico system [5].
- Inferred mineral resource of 1.6 million tonnes, grading 0.32% Cu, 0.20% Ni, and 1.7 g/t Pt+Pd [7].



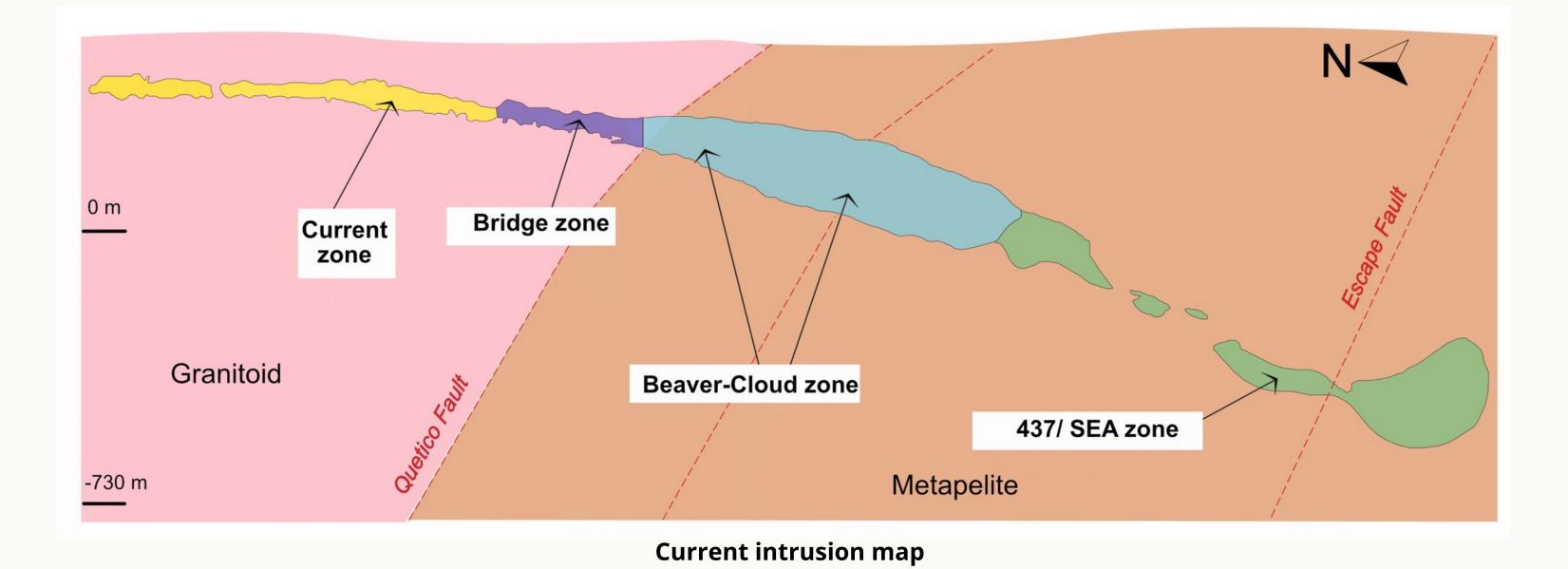
[1] Bleeker et al. (2020) [5] Williams (1991)

[6] Modified after Nordmin Engineering Ltd (2021)

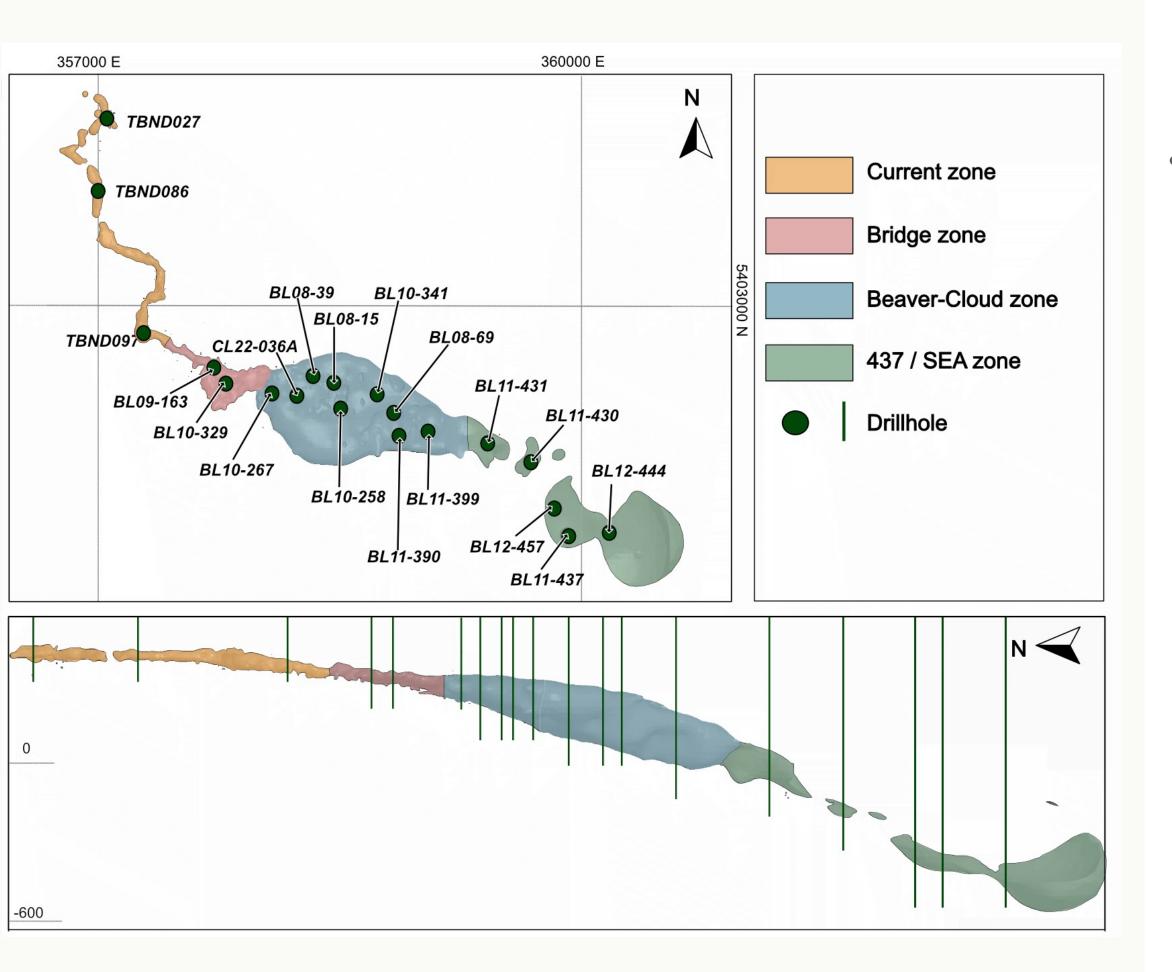
Current intrusion map [6]

Current Intrusion

• Current intrusive is divided into: Current zone, Bridge zone, Beaver-Cloud zone, and **437 SEA zone** [8].



[8] Brzozowski et al. (2023)

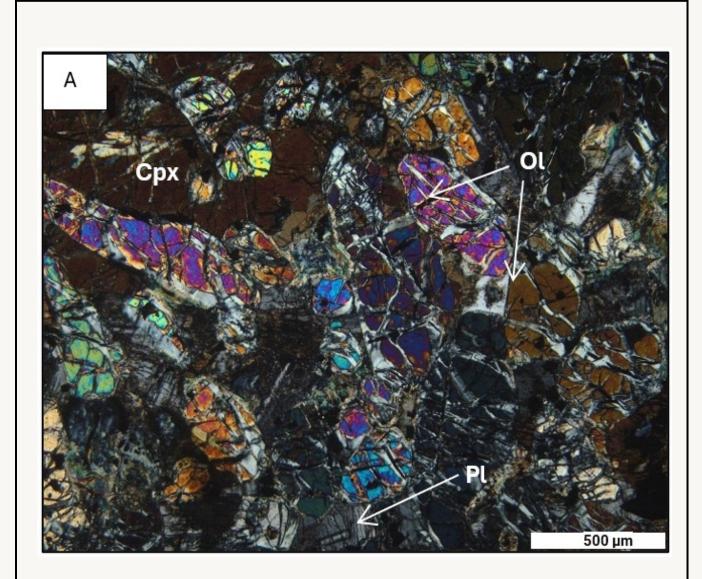




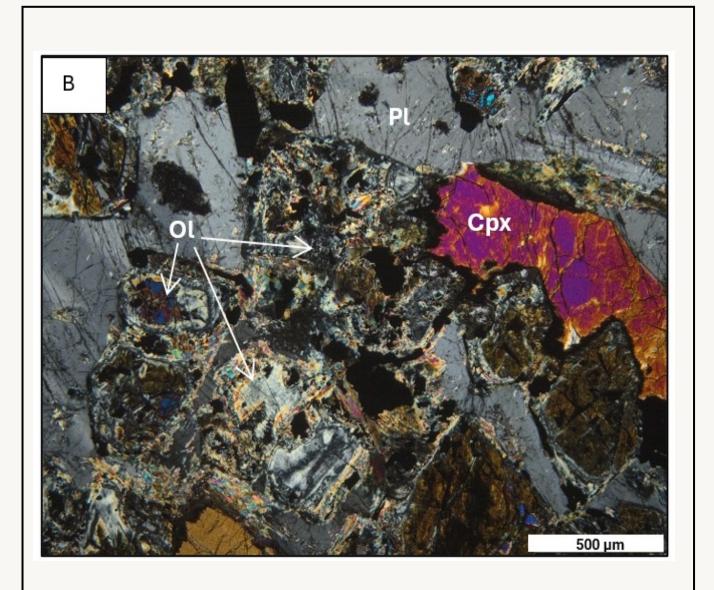
Nineteen selected drillholes:

- Thirty-six samples for petrographic examination and SEM-EDS.
- → Fifty-three samples for geochemical studies.
- Thirty-eight samples for stable isotope analysis.
- → Ten samples for radiogenic neodymium and strontium isotope analysis.

Petrography



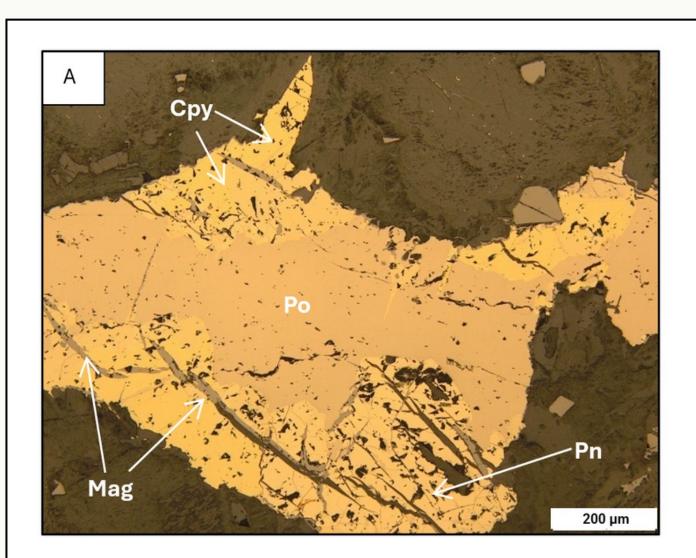
• Fine-medium grained wehrlite and Iherzolite: 40-70% olivine, 0-3% orthopyroxene, 5-25% clinopyroxene, and 5-15% plagioclase.



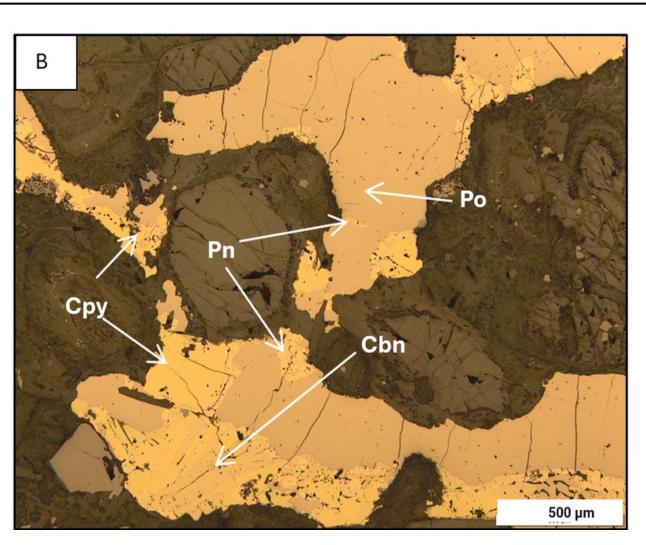
• Fine- medium grained olivine gabbronorite ± troctolite: 40-55% plagioclase, 35-45% olivine, and 5-10% pyroxene.

Petrography

Sulfides 0-9%, averaging 50% pyrrhotite, 40% chalcopyrite, and 9% pentlandite, 1% cubanite.



Fine-to-medium-grained pyrrhotite, fine-grained chalcopyrite, and fine-grained pentlandite. Pentlandite and chalcopyrite are weakly altered by magnetite.



Medium-grained pyrrhotite with very fine-grained pentlandite exsolutions. Medium-grained chalcopyrite with fine-grained cubanite exsolutions.



Paragenetic Sequence

Magmatic Phase

SILICATES			Temperature decreases
1. Magmatic	A. Antigorite	2. hydrothermal alteration B. Lizardite-chrysotile	n C Talc - carbonate
	Act-Tr Atg Clc Ep	Z+ctl	
SULFIDES-OXIDES			
1. Mag	gmatic	2. hydrotherm A-B Py-mag-chm	A-B Secondary Po
ISS	Cbn Cpy Po	Mag	

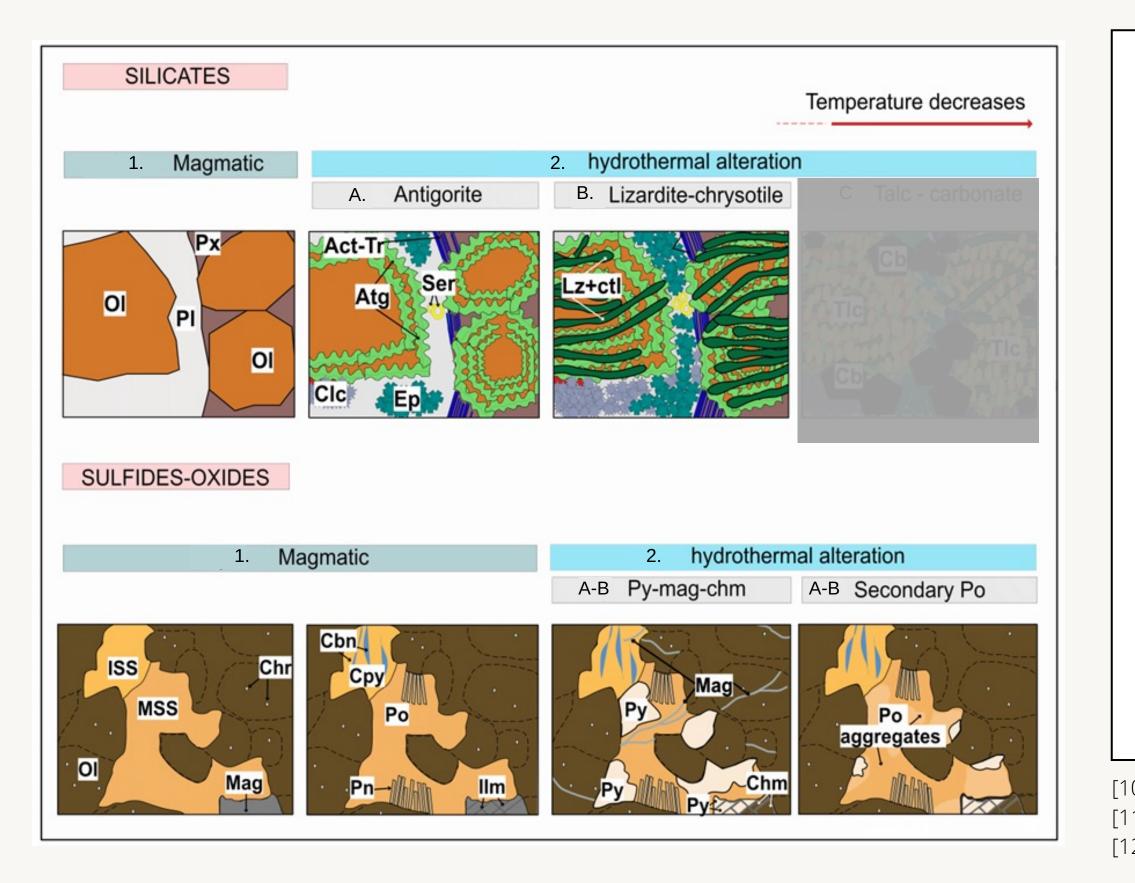
• Early formation of olivine phenocrysts enclosing Cr-spinel crystals, suggesting that the chromite crystallized before the olivine.

• As the temperature of the magmatic system decreased, oikocrysts of pyroxene enclosed the olivine crystals, with subhedral **plagioclase** subsequently filling the remaining interstices.

• Oxides and sulfides are interstitial to the mafic silicates, indicating **crystallization** subsequent to the olivine and pyroxene

Paragenetic Sequence

Hydrothermal Phase



 Domain A is characterized by antigorite, actinolite-tremolite, clinochlore, epidote, sericite, pyrite, millerite, secondary pyrrhotite, chamosite, and secondary magnetite. Formed by
 >300°C fluids ; [10,11,12]) with acidic pH [13, 14].

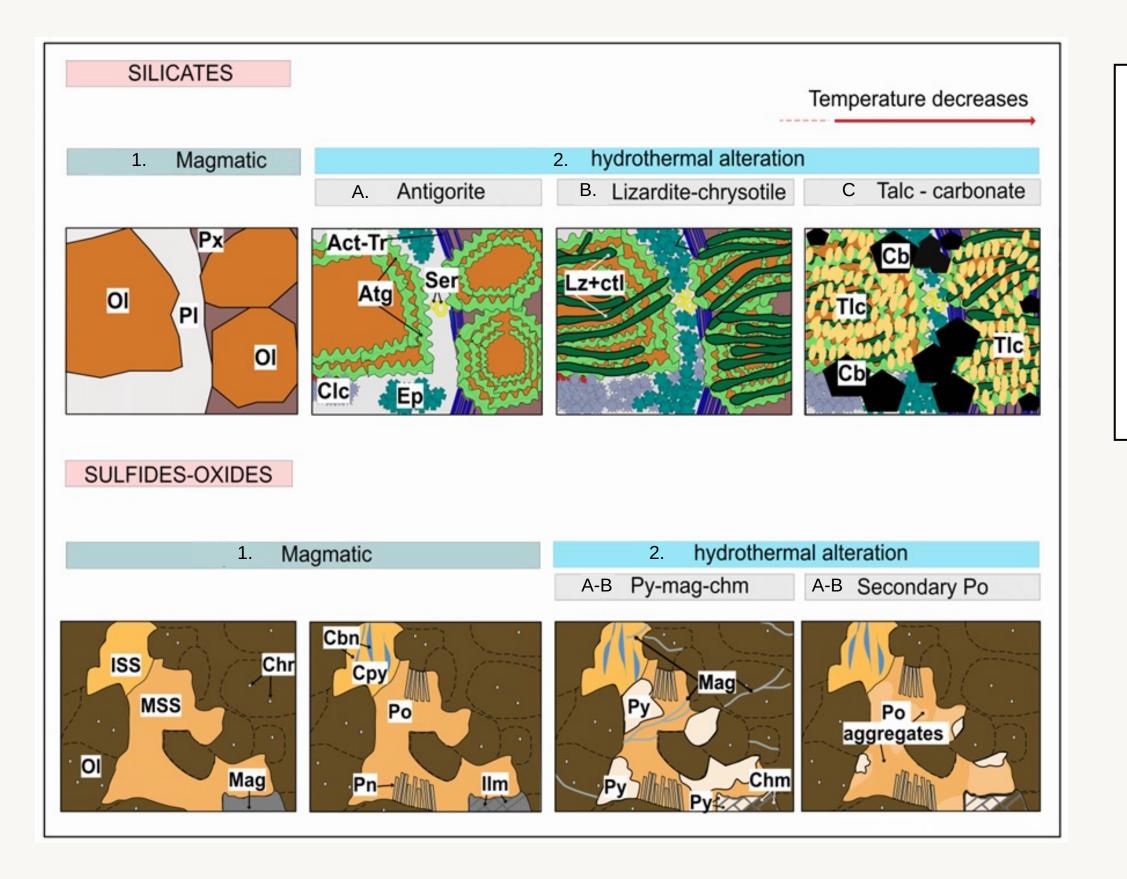
• Domain B consists of lizardite-chrysotile and an increase in the modal abundances of clinochlore, epidote, sericite, pyrite, millerite, and secondary magnetite compared to Domain A. Formed by <300°C fluids [10,11,12) with acidic pH [13, 14].

[10] Evans (2004)[11] Lagat (2009)[12] Welch and Marshall (2015)

[13] O'Hanley (1996) [14] Fulignati (2020)

Paragenetic Sequence

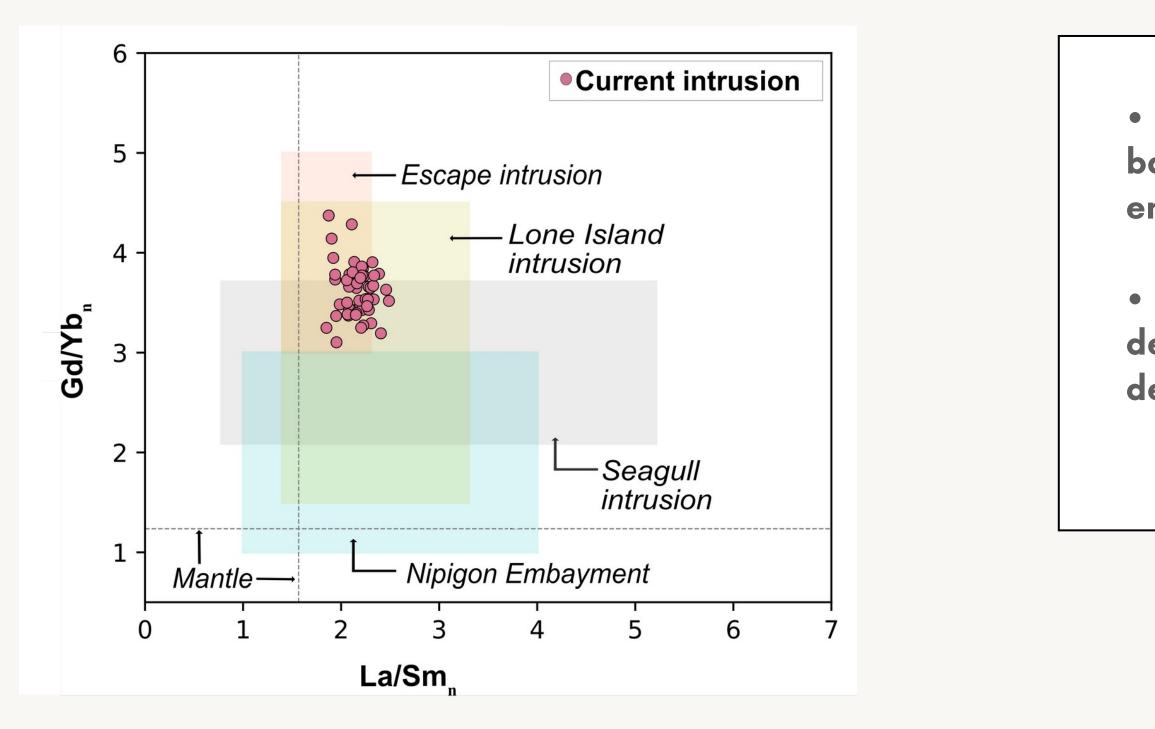
Hydrothermal Phase



Domain C comprises talc and carbonate minerals that have replaced the secondary minerals present in Domains A and B. Associated with later fluids below 50°C [15,16], crystallizing under basic pH conditions [17].

[15] Barnes et al. (1973)[16] Kelemen and Matter (2008)[17] Chaliulina (2019)

Magmatic Signature Whole-Rock Geochemistry



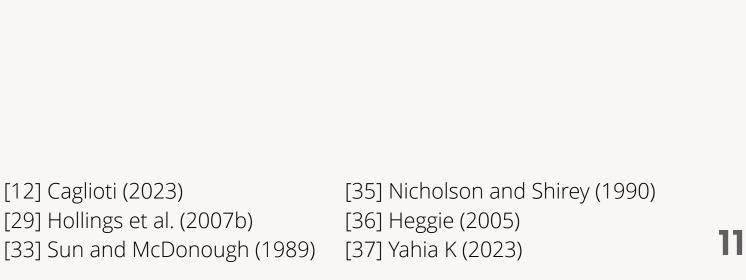
Chondrite-normalized, whole-rock-trace variation for Current intrusion. Nipigon Embayment, Escape, Lone Island, and Seagull intrusion [12,29,35,36,37]. Mantle values and normalizing values from [33].

[12] Caglioti (2023)

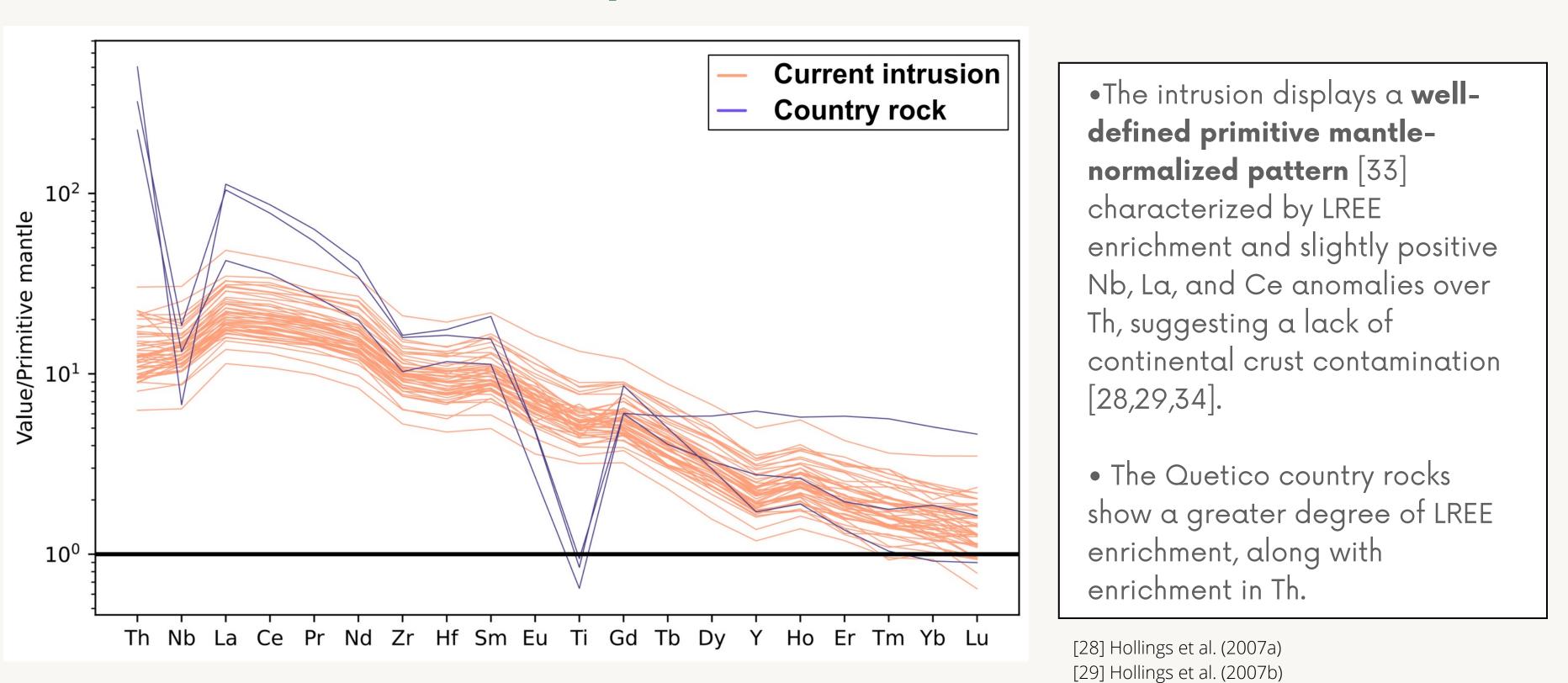


The **La/Smn** values indicate a basaltic magma derived from an enriched mantle plume [35].

• The Gd/Ybn values suggest a lower degree of partial melting at greater depths.



Magmatic Signature Whole-Rock Geochemistry

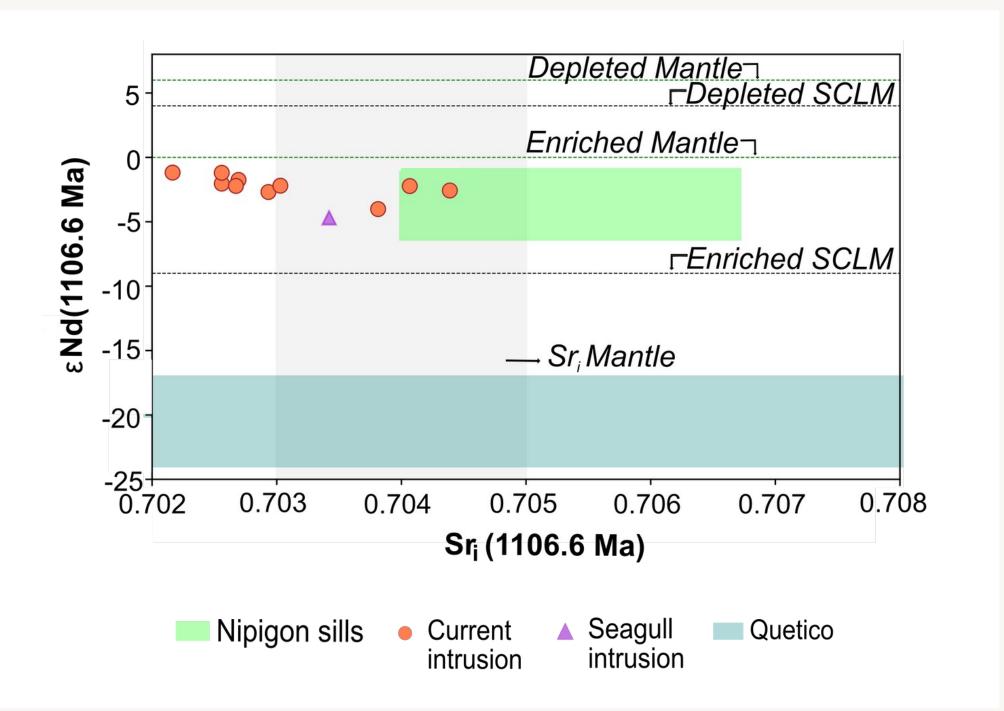


Primitive mantle-normalized spider diagram for Current intrusion and Quetico country rocks. Normalizing values from [33].



- [33] Sun and McDonough (1989)
- [34] Hofmann (1997)

Magmatic Signature Radiogenic Isotopes



Mantle, SCLM, and Quetico country rock values were taken from [35,38,39,40,41,42] and the Coubran volcanics, Nipigon sills, and Seagull intrusion data were taken from [29,31,36].



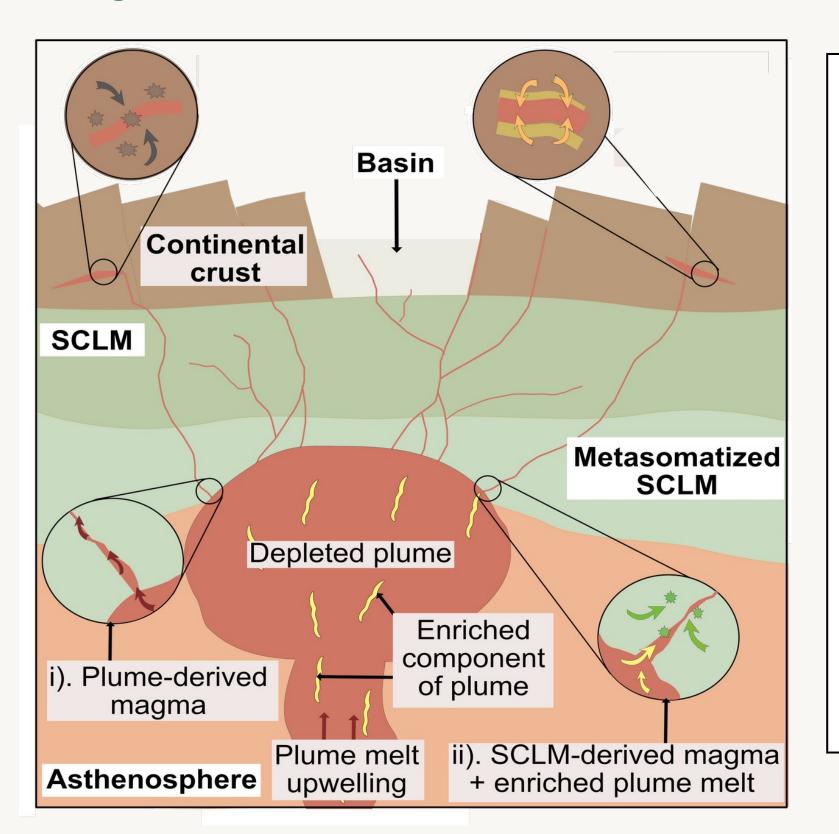
• The intrusion exhibits slightly lower ENd values compared to the typical values of the mantle source at 1100 Ma [35,38].

• The absence of geochemical anomalies indicating crustal contamination suggests that an **enriched** Subcontinental Lithospheric Mantle (SCLM) likely interacted with the **parental magma**.

[29] Hollings et al. (2007a) [31] Cundari R (2012)

- [35] Nicholson and Shirey (1990)
- [36] Heggie (2005)
- [38] Hergt et al. (1989)
- [39] Shirey et al. (1994) [40] Shirey (1997) [41] Henry et al. (1998) [42] Pan et al. (1999) [43] Rooney et al. (2022)

Magmatic Signature Magmatic Evolution



i) depleted plume-derived magma + crustal **contamination** (melting of xenoliths and/or boundaries of the crustal rock; [8]), scenario proposed for the Coubran basalts and Nipigon sills by [33].

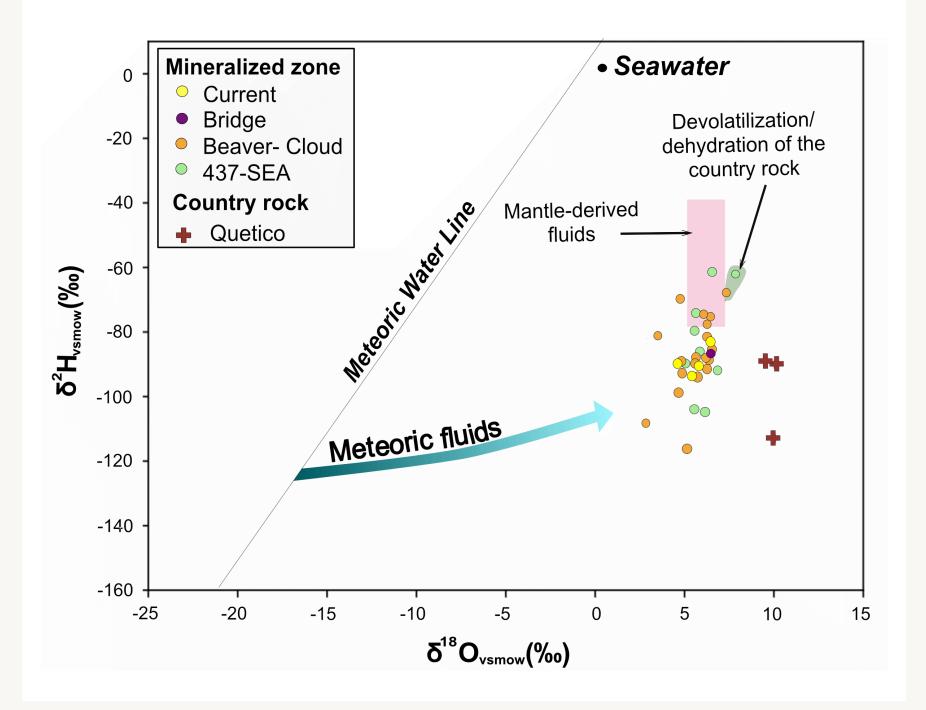
ii) SCLM-derived magma + enriched plume melt + sulfide saturation by thermal aureole in the crust (thermal decomposition of sulfides present in the wall rock [8]), scenario proposed for the Current intrusion.

Schematic model of the Keweenawan plume interaction on its way to the crust.



• During the **transportation** of the melt from its source in the mantle through the lithosphere to the base of the crust, varying levels of contamination may occur for some magmas associated with the Keweenawan plume:

Hydrothermal Signature Stable Isotopes



Stable isotopes values of bulk rock of the Current intrusion and the surrounding country rock.

- [55,56]).

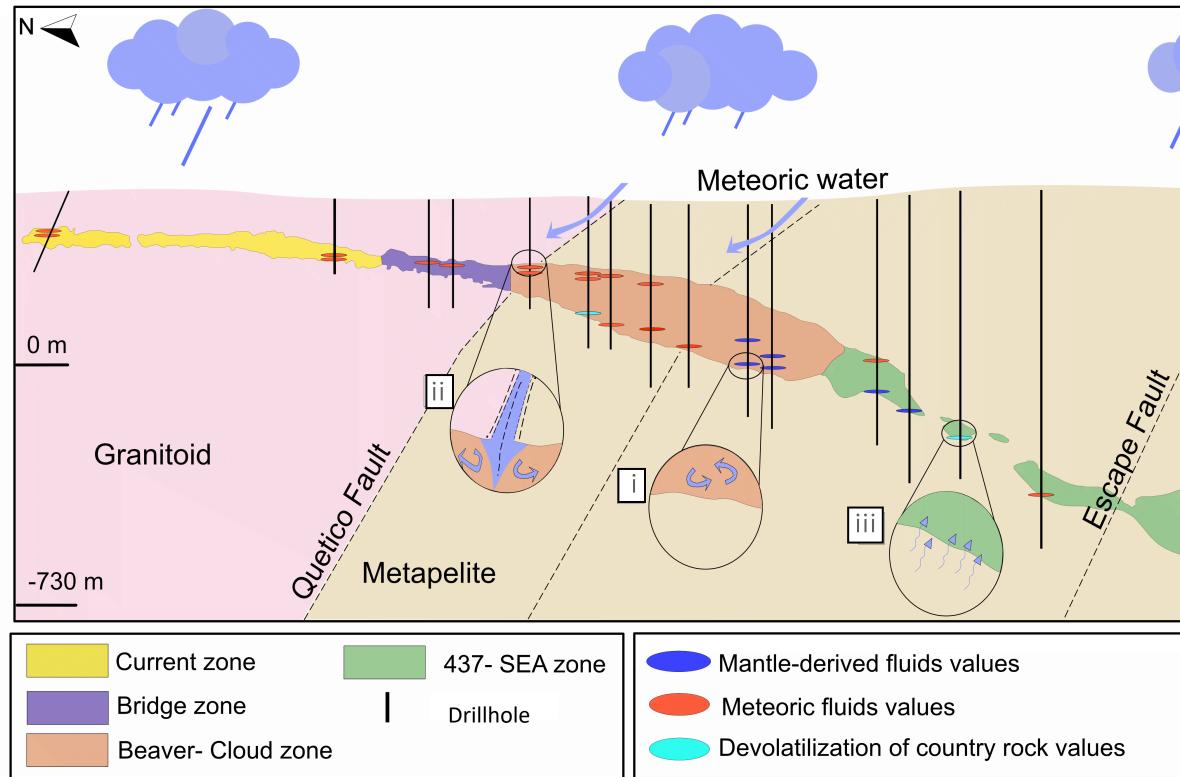
[55] Ripley et al. (1993) [47] Taylor H (1968) [51] Loewen et al. (2019) [56] Park and Ripley (1999) [48] Kyser et al. (1982) [52] Moine et al. (2020) [49] Mattey et al. (1994) [53] Bindeman et al. (2022) [50] Eiler (2001) [54] Ripley and Al-Jassar (1987

• i) mantle-derived fluids (δ^2 H from -40 to -80‰, δ¹°O from 5.5 to 7.0‰ [47, 48,49,50,51,52,53].

• ii) meteoric fluids ($\delta^2 H < -80\%$, $\delta^{18}O < 5.5\%$;

• iii) crustal-derived fluids by devolatilization/ dehydration crustal-derived fluids by devolatilization / dehydration ($\delta^{18}O > 7\%$; [54]).

Hydrothermal Evolution

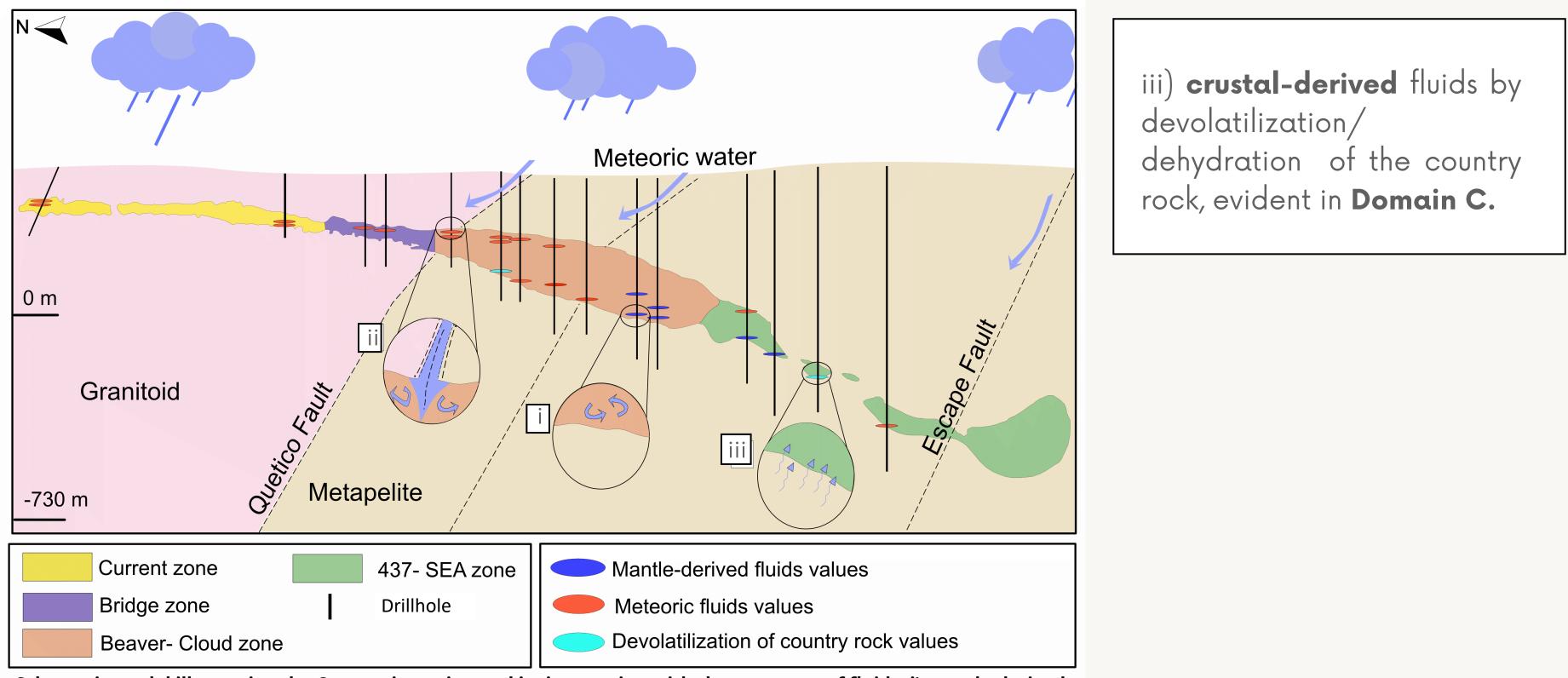


Schematic model illustrating the Current intrusion and its interaction with three sources of fluids: i) mantle-derived fluids; ii) meteoric fluids; iii) crustal-derived fluids.

i) magmatic fluids
 preserved chiefly at the
 bottom of the intrusion,
 generated the secondary
 mineralogical assemblages
 of Domain A and B.

ii) meteoric fluids recorded mostly at the top of the intrusion, generated a secondary mineralogy similar to the derived from the magmatic fluids.

Hydrothermal Evolution

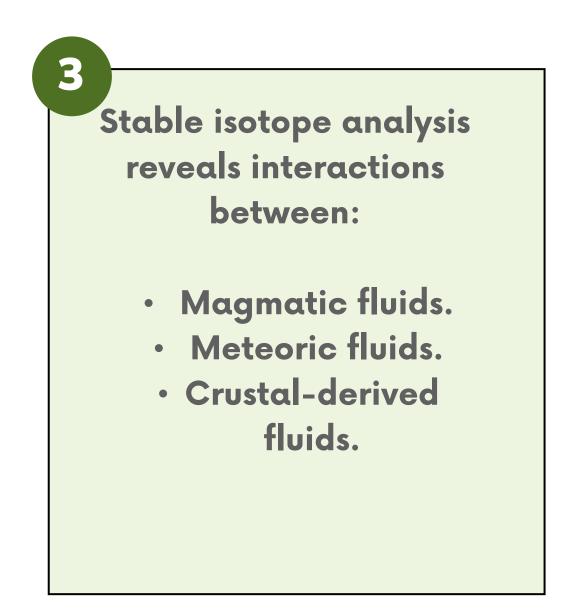


Schematic model illustrating the Current intrusion and its interaction with three sources of fluids: i) mantle-derived fluids; ii) meteoric fluids; iii) crustal-derived fluids.

Conclusions

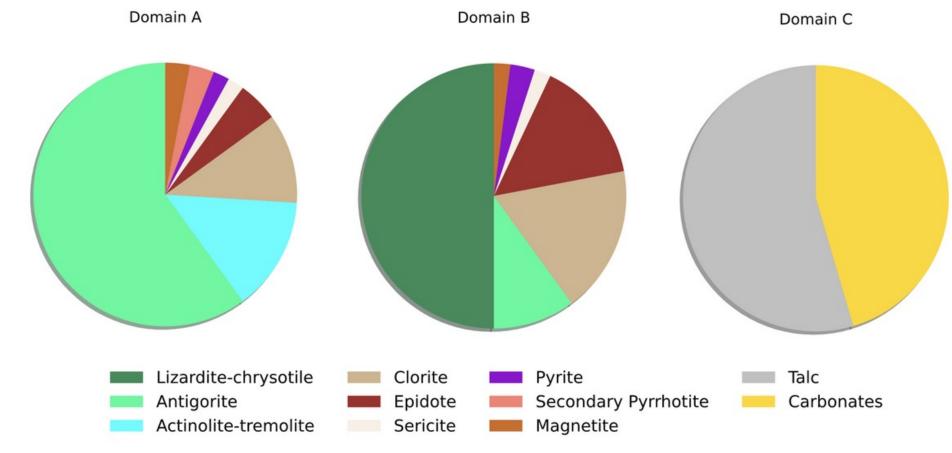
The Current deposit has a primitive mantlenormalized pattern and enriched nature of the magma suggests a basaltic magma derived from an enriched mantle plume with low partial melting at greater depths with no continental crust contamination.

The slightly lower ENd values indicate interaction with an enriched SCLM.



Conclusions

	Secondary silicates						ndary ides	Secondary oxides			
Mineral domains	Antigorite	Lizardite- chrysotile	Chlorite	Epidote	Actinolite- tremolite	Talc	Sericite	Pyrite	Pyrrhotite	Magnetite	Carbonates
Domain A	+		+	+	+		+	+	+	+	
Domain B	+	+	+	+			+	+		+	
Domain C						+					+



The intrusion exhibits three distinct alteration domains:







Thank you